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Oceanographic Conditions on the Flemish Cap in NAFO Division 3M During the Summer of 2004

by

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ABSTRACT

Oceanographic data from the summer of 2004 on the Flemish Cap are examined and compared to the long-term (1971-2000) average. The cold near-surface temperatures (0.5° to 2°C below normal) experienced over the Cap from 1993 to 1996 had warmed to 0.5°-1.5°C above normal by the summer of 1997, which increased further to 2°C above normal by the summer of 1999. Upper layer temperatures over the Flemish Cap during the spring of 2001 and the summer of 2002 generally showed a downward trend with temperatures decreasing to below normal values. During the summer of 2003, temperatures directly over the Cap were highly variable while adjacent areas showed significant positive anomalies and during 2004 they increased to above normal values. Near bottom temperatures over the Cap during 2004 were $>4^{\circ}$ C, which was above normal by 0.5°C over the shallow areas of the Cap. Salinities over most of the upper water column during the summer of 2002 to 2004 were generally saltier-than-normal (0.25-0.5). In the deeper water (>100 m depth) salinities were about normal. In general, the colder-than-normal temperatures experienced over the continental shelf and on the Flemish Cap from the late-1980s up to the mid-1990s moderated by the summer of 1996 and continued to warm until 1999. During the summer of 2000 and into the spring of 2001 the observations indicate a reversal in the recent warm trend in some areas with water column temperatures decreasing to near normal values. However during the summers of 2003 and 2004 most areas of the water column again experienced an increase in both temperature and salinity. During 2004 and throughout most of the 1990s and early 2000s summer chlorophyll levels in the upper 100-m of the water column over the Cap were higher compared to the adjacent Grand Bank. Dissolved oxygen levels were about normal for the region with super-saturated values in the near surface layers reaching 105%. The circulation pattern during the summer of 2004 was dominated by the southward flowing Labrador Current to the east of the Cap indicating a reduced gyre circulation.

1. Introduction

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The Flemish Cap is an isolated bank located east of the Grand Bank of Newfoundland centred at about 47° N, 45° W with minimum water depths of 126-m (Fig. 1 top panel). To the west, the Flemish Pass with maximum water depths of about 1100 m separates the Cap from the Grand Bank. The water mass over the Flemish Cap is a mixture of Labrador Current and North Atlantic Current water, the general circulation of which is shown in Fig. 1 (bottom panel). Since the early-1970s the water masses over the Atlantic continental shelves have been dominated by three anomalously cold periods: early-1970s, mid-1980s and the early-1990s (Colbourne *et al.*, 1994). During these periods strong positive winter North Atlantic Oscillation (NAO) index anomalies were mainly responsible for colder than normal air temperatures over the Northwest Atlantic resulting in increased ice cover and eventually colder and fresher than normal oceanographic conditions over the continental shelf of Atlantic Canada (Drinkwater, 1996). An examination of the time series of historical temperature data from the Flemish Cap indicated similar conditions existed on the Flemish Cap. During the mid-1990s however, the ocean climate in this region returned to more normal like conditions and by the late-1990s ocean temperatures in many areas reached record highs (Colbourne, 1999, 2000; Colbourne *et al.*, 2004).

Keeley (1981) published monthly mean temperature and salinity data from 1910-1980 along the standard Flemish Cap Section (Fig. 1). Drinkwater and Trites (1986) published spatially averaged temperature and salinity from all available bottle data from 1910 to 1982 over the Flemish Cap area. Numerous reviews and studies of the physical oceanography around the Flemish Cap were conducted during the Flemish Cap Project of the late-1970s and early-1980s (Hays *et al.*, 1978; Bailey, 1982; Akenhead, 1981). More recent reviews of oceanographic conditions in the region compared 1993 and 1995-2003 observations with the long-term mean and also presented time series of temperature and salinity anomalies at various depths around the Cap (Colbourne, 1993, 1995, 1996, 1997, 1998a, 1998b, 1999, 2000, 2001, 2002, 2003). Stein (1996) summarised the most recent studies of oceanographic conditions on the Flemish Cap. Additionally, Cervifio and Prego (1996), Garabana *et al.* (2000), Lopez (2001) and Cabanas (2002, 2003) have presented hydrographic conditions on the Flemish Cap in July from fisheries research surveys conducted by the European Union. This manuscript presents an update to these studies by including data up to the summer of 2004 and represents the 12th such review of oceanographic conditions on and around the Flemish Cap in support of the annual shrimp assessment for NAFO Division 3M.

This report describes oceanographic variability on the Flemish Cap during the summer of 2004 with a comparison to the long-term mean, based on all available historical data. The normal has been defined as the 30-year period from 1971-2000. The 2004 observations were made by Canada's Department of Fisheries and Oceans oceanographic survey during the summer of 2004. During these surveys oceanographic observations were made along the standard NAFO Flemish Cap section at 47°N latitude (Fig. 1, top panel). Physical oceanographic measurements included vertical profiles of temperature, salinity, chlorophyll, dissolved oxygen and ocean currents.

2. Average Temperature and Salinity

Vertical distributions of the temperature and salinity fields over the Flemish Cap along 47°N averaged over the period of July 10 to August 7 based on all available historical data from 1971-2000 are shown in Fig. 2. These dates were chosen to span a one-month period centred on the 2004 observations. No other attempts were made to adjust the mean for possible temporal biasing arising from variations in the number of observations within the time interval.

The average near-surface temperature across the Flemish Cap for this period ranges from $9^{\circ}-11^{\circ}$ C (Fig. 2, upper panel). Below the thermocline (approximately 50 m) temperatures range from $2^{\circ}-4^{\circ}$ C on the Flemish Pass side of the Cap to $3^{\circ}-4^{\circ}$ C over the Cap and to the east of the Cap. Water temperatures at depths >200 m in the Flemish Pass, and on the eastern slopes, range from $3.5^{\circ}-4^{\circ}$ C. Bottom temperatures over the Cap and on the eastern slopes of the Cap also range from $3.5^{\circ}-4^{\circ}$ C. The corresponding average surface salinities (Fig. 2, bottom panel) range from <33.25 in the Flemish Pass to about 33.5-33.75 over the Cap and eastward of the Cap. Near bottom over the Cap, in water depths of 150 to 300 m, salinities range from 34.5 to 34.75. In general, cold subsurface and relatively fresh surface water from the Newfoundland Shelf normally influences the water mass characteristics over most of the Flemish Cap.

3. 2004 Temperature and Salinity

Surface temperatures over the Flemish Cap during the summer of 2004 (Fig. 3 top panel) ranged from about 10° - 12° C, which were generally above normal in most areas, although there were isolated areas with below normal values. Below the surface layer, temperatures of the intermediate depth waters were above normal across the region by up to 2°C. The edge of the cold intermediate layer (CIL) waters with temperatures <3°C from the Newfoundland Shelf were present in the Flemish Pass during July of 2004. Near-bottom temperatures over the Cap in water depths from 150-175 m were >4°C, which was above normal by more than 0.5°C. Below 200 m depth bottom temperatures were about normal during the summer of 2004.

Near-surface salinities observed during the summer of 2004 (Fig. 4 top panel) show values ranging from <33.5 over the Flemish Pass to between 34-34.25 over the Cap. Salinities near bottom over the Cap in water depths around 150 m were generally >34.5 and >34.75 below 250 m depth. In general these values were above the long-term mean in the upper 75 m of the water column by up to 0.5 in some areas. Below 100 m depth and near bottom salinities were about normal ranging from 34.5 to 34.8.

4. Long-Term Trends in Temperature and Salinity

The time series of annual temperature anomalies in the upper water column on the Flemish Cap (Fig. 5 left panels) show considerable variability with cold periods during most of the 1970s, mid-1980s and the late-1980s to the mid-1990s. During most of the 1960s, late-1970s to early-1980s and again during the late-1990s temperatures were generally warmer-than-normal. The cold period beginning around 1971 continued until about 1977. From 1978 to 1984 temperature anomalies showed a high degree of variability in the upper water column with a tendency towards positive anomalies. By 1985 in the top 100-m of the water column negative temperature anomalies had returned. This cold period moderated briefly in 1987-88 but returned again by 1989 and continued until the mid-1990s. From 1995 to 1998 temperatures moderated and remained above normal during the summer of 1999 in the depth range from 0 to 150-m. During 2000 and 2001 temperatures decreased significantly over 1999 values but remained near normal at 150 m and slightly above normal at shallower depths. Temperatures at all depth increased during 2003 and 2004 to above normal values by >1°C at the surface and >0.5°C at 150 m depth. In general, temperature trends on the Flemish Cap are very similar to those observed on the Newfoundland Shelf during the same period (Colbourne and Foote, 2000).

The time series of annual salinity anomalies (Fig. 5 right panels) show fresher-than-normal conditions during the mid-1970s and mid-1980s in the upper 100-m of the water column, with peak amplitudes reaching 0.9 below normal. In general, the magnitude of the salinity anomalies decreases with increasing depth. The trend in salinity values during the early-1990s was mostly below normal until 1995 from the surface to about 100 m depth. From 1996 to 1998 salinities were generally above normal. During 1999-2000 salinities were near the long-term average on the Flemish Cap but increased to above normal values during 2001-2002. Upper layer salinities during 2003 increased to some of the highest values observed in the time series. During 2004 they decreased slightly over 2003 values but remained above the long term mean for the 4th consecutive year. In general, salinity anomalies are very similar to those at Station 27 and elsewhere on the continental shelf over similar depth ranges (Colbourne 1998b). It should be noted however, that unlike the time series of anomalies from fixed points (e.g. Station 27), these anomalies, both temperature and salinity, are based on data collected over a larger geographical area and therefore may exhibit variability due to spatial differences in the monthly estimates. In addition, the annual values may be based on only a few monthly estimates for the year. Therefore caution should be used when interpreting short time scale features of these series. The long-term trends however, generally show real features.

5. 2004 Chlorophyll and Dissolved Oxygen

The vertical distribution of dissolved oxygen saturation during the summer of 2004 along the standard NAFO section across the Grand Bank and the Flemish Cap is shown in Fig. 6. These data were collected in conjunction with the temperature and salinity data using a SBE-43 type polarographic element dissolved oxygen sensor interfaced to a Seabird-911 CTD system. The oxygen sensor was factory calibrated at zero and air-saturated water oxygen levels and also field calibrated by taking water samples at standard depths. The oxygen levels of the samples were determined by semi-automated analytical chemistry using a modified Winkler titration technique. The sensor readings were then corrected by using least-squares fit (slope 0.98, intercept 0.32) of the titration measurements to the electronic sensor measurements. Over the Flemish Cap dissolved oxygen saturation levels during the summer of 2004 ranged from 100-105% from the surface to about 50-m depth (Fig. 6). From 50-100 m depth, saturation values decreased from 100% to 90% and near bottom values were generally around 75-85% saturation. In general oxygen levels observed along this section are similar to that observed during most of the 1990s and are typical of the highly oxygenated water column of this region.

The vertical distribution of relative chlorophyll concentrations for the summer of 2004 across the Grand Bank and the Flemish Cap is shown in Fig. 7. These data were collected in conjunction with the temperature and salinity data using a fluorometer interfaced to a Seabird-911 CTD system. No field calibrations were applied to the chlorophyll values presented here. The chlorophyll concentrations were generally low over the Grand Bank but somewhat higher over the Flemish Cap. Maximum chlorophyll values over the Cap were confined to the surface layer of about 80-m thick (Fig. 7). The higher chlorophyll values over the Flemish Cap appear to extend into mid-summer and may indicate an extended offshore plankton bloom relative to the Newfoundland Shelf areas (Colbourne 2000).

6. Circulation

The general circulation in the vicinity of the Flemish Cap consists of the offshore branch of the Labrador Current which flows through the Flemish Pass on the Grand Bank side of the continental slope and a jet that flows to the east, north of the Cap, which then flows southward. To the south, the Gulf Stream flows to the northeast merging with the Labrador Current to form the North Atlantic Current (Fig. 1 bottom panel). In the absence of strong wind forcing (mainly summer) the circulation over the Flemish Cap is dominated by a topographically induced anticyclonic gyre over the central portion of the bank (Kudlo *et al.*, 1984; Ross, 1981). The stability of this circulation pattern may influence the retention of ichthyoplankton on the bank and is probably a factor in determining the year-class strength of various fish and invertebrate species, such as cod, redfish and shrimp (Kudlo and Borovkov, 1977; Kudlo and Boytsov, 1979).

The anticyclonic motion of the water mass around the Flemish Cap was first described by Kudlo and Burmakin (1972), Kudlo and Borovkov (1975) and Kudlo *et al.* (1984) using geostrophic currents estimated from density measurements. Gil *et al.* (2002) and Colbourne and Foote (2000) provided additional analysis based on more recent data. The geostrophic currents perpendicular to the 47° N section calculated from the density data collected during the summer of 2003 and 2004 are shown in Fig. 8. These estimates, which are referenced to 300-m, or the bottom, in water depths <300 m, show some of the well-known features of the circulation. The strong baroclinic component of the offshore branch of the Labrador Current west of the Flemish Pass and the northward flowing water of the North Atlantic Current east of the Cap are evident. The current estimates over the Cap itself show evidence of anticyclonic circulation with northward flowing water to the western slopes of the Cap and the other directly over the Cap. The northward component of the gyre on the western side of the Cap during 2004 was similar to that observed in 2002, much weaker than that observed during the summers of 1999, 2000 and 2003 (Colbourne, 2000, 2003). Finally, during the summer of 2004, it appears that the circulation pattern over the Cap was dominated by the southward flowing Labrador Current to the east of the Cap indicating a reduced gyre circulation.

7. Summary

The cold near-surface temperatures (0.5° to 2°C below normal) experienced over the Cap from 1993-1996 had warmed to 0.5°-1.5°C above normal by the summer of 1997, which increased further to 2°C above normal by the summer of 1999. Upper layer temperatures over the Flemish Cap during the spring of 2001 and the summer of 2002 generally showed a downward trend with temperatures decreasing to below normal values. During the summer of 2003, temperatures directly over the Cap were highly variable while adjacent areas showed significant positive anomalies and during 2004 they increased to above normal values. Near bottom temperatures over the Cap during 2004 were $>4^{\circ}$ C, which was above normal particularly over the shallow areas of the Cap. Salinities over most of the upper water column during the summer of 2002 to 2004 were generally saltier-thannormal (0.25-0.5). In the deeper water (>100 m depth) salinities were about normal. In general the colder than normal temperatures experienced over the continental shelf and on the Flemish Cap from the late-1980s up to the mid-1990s moderated by the summer of 1996 and continued to warm until 1999. During the summer of 2000 and into the spring of 2001 the observations indicate a reversal in the recent warm trend in some areas of the water column with temperatures decreasing to near normal values in most areas. During the summers of 2003 and 2004 however, most areas of the water column experienced an increase in both temperature and salinity. During 2004 and throughout most of the 1990s and early-2000s summer chlorophyll levels in the upper 100-m of the water column over the Cap were higher compared to the adjacent Grand Bank. Dissolved oxygen levels were about normal for the region with super-saturated values in the near surface layers reaching 105%. In general oxygen levels observed along this section are similar to that observed during most of the 1990s and are typical of the highly oxygenated water column of this region. The current estimates over the Cap indicate that the circulation pattern was dominated by the southward flowing Labrador Current to the east of the Cap indicating a reduced gyre circulation during the summer of 2004. During most oceanographic surveys of the 1990s and up to 2004 ocean currents around the Flemish Cap area were measured with hull-mounted acoustic Doppler current profilers (ADCPs). In general, the details of the circulation patterns measured with ADCPs differ significantly from the geostrophic estimates, thus showing the potential importance of wind driven and tidal currents, however, they both show similar features in the overall circulation pattern.

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Fig. 1. Areal map showing the standard Flemish Cap section in NAFO Sub-area 3 and the Flemish Cap region in NAFO Div. 3M (top panel) and the major circulation features around the Flemish Cap area (bottom panel).



Fig. 2. Average temperature (in °C) and salinity on the Flemish Cap based on all data collected between July 9 to August 6 for the years 1971-2000.



Fig. 3. The vertical distribution of temperature and temperature anomalies (in °C) over the Flemish Cap (along 47°N) for the summer of 2004.



Fig. 4. The vertical distribution of salinity and salinity anomalies over the Flemish Cap (along 47° N) for the summer of 2004.



Fig. 5. Annual temperature and salinity anomalies at standard depths on the Flemish Cap in NAFO Division 3M. The solid line represents a 5-year running mean.



Fig. 6. The vertical distribution of dissolved oxygen percent saturation along 47°N for the summer of 2004.



Fig. 7. The vertical distribution of chlorophyll concentrations (mg/l) along the 47°N transect for the summer of 2004.



Fig. 8. The vertical distribution of the N-S geostrophic current field (in cm/s) over the Flemish Cap during the summer of 2003 (top panel) and 2004 (bottom panel) estimated from the density data. Negative currents (blue) are southward and positive currents (red) are northward.